# Site Investigations for Natural Terrain Hazard Assessments

Steve Parry (GEO)

# Talk is divided into two parts

 use of geomorphology and API to develop hazard models

• Focused detailed GI to assess hazard models

- Site Investigation does not equal Ground Investigation
- For Natural Terrain a holistic approach is required i.e. all encompassing
- Given difficult access, dense vegetation, the size of site any SI needs to be focused and this is best achieved by the development of models
- To develop models you be able to "read" the landscape which requires skills and experience, these including an understanding of engineering geology, geomorphology and natural terrain hazards. However, there is limited experience of geomorphological mapping in Hong Kong and even less experience with respect to developing models of landscape evolution

# Geomorphology

"is the study of the forms of the surface of the earth, their origin, the processes involved in their development, the properties of the materials of which they are made and predictions about their future form, behavior and status"

(D Brunsden, 5th Glossop Lecture, 2002)

"While engineering geological mapping is perhaps concerned with a snapshot in time regarding the properties of materials and their immediate or short term engineering implication, geomorphological mapping, in theory, takes in a greater sweep of time, combining the recent geological past with the present geomorphology and its foreseeable future. Deriving an engineering evaluation from an integration of these two approaches combines the short term static with the longer term dynamism of the landscape: an evaluation which is critical to the survival of engineering structures throughout and beyond their notional design life"

Hearn, (2002) Engineering geomorphology for road design in unstable mountainous areas: lessons learnt after 25 years in Nepal. QJEG Vol 35 p143-154

- Need to understand how the landscape evolved and what that can tell us about future hazards
- Key component is API Interpretation
- Not <u>observation</u> which is what many reports contain often being simple site histories this is of little use for NTHA.
- Key photographs are 1963/4 use both high and low level as these help put the site in <u>context</u> i.e. where it is within the landscape
- In addition colour photographs, particular early ones are beneficial for vegetation changes, rock outcrops etc.
- A site reconnaissance is also critical during the API helps you to confirm both your observations and interpretation

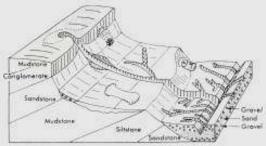
## **Common Problems:**

- Geomorphological boundaries derived from API are not verified in the field
- Assessments are often restricted to site boundaries. Landforms do not evolve in isolation and in order to develop a model a sufficient large area has to be studied
- When geomorphological mapping is carried out it commonly shows only rudimentary morphological data
- Single mapping approaches are often used e.g. regolith mapping. However as all sites are to some degree unique, some approaches are more useful than others (see TGN 22)
- Studies do not focus on the hazard and risk

# Brief introduction to geomorphological approaches

For more details refer to TGN 22

#### WORKING PARTY REPORT



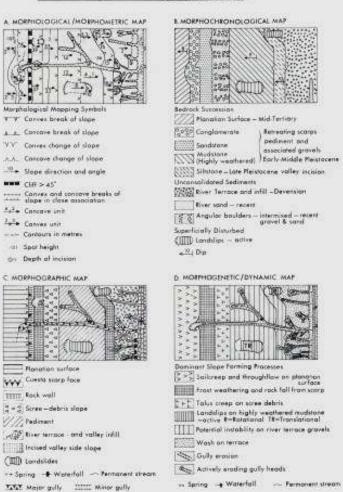
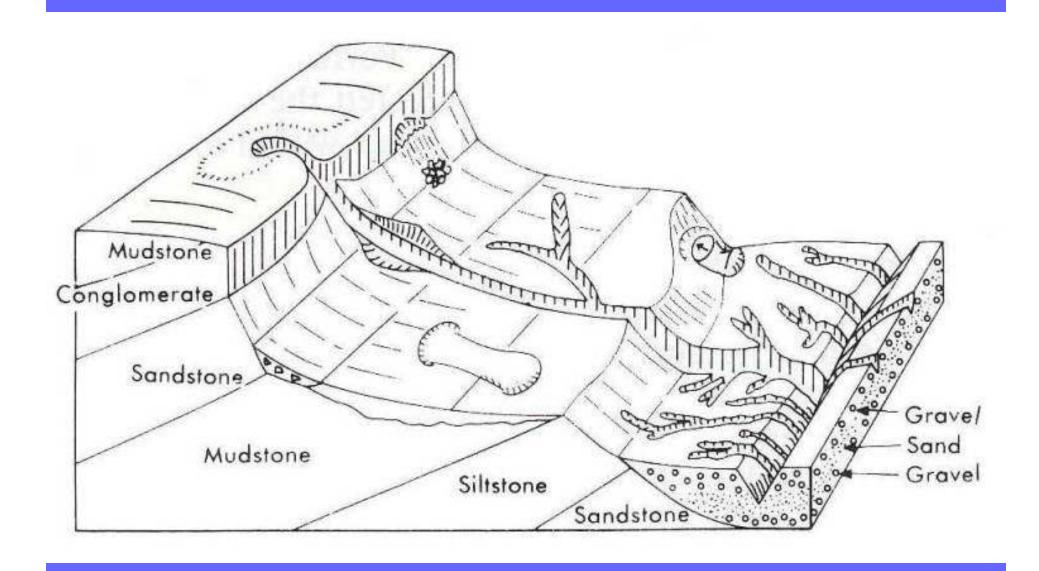
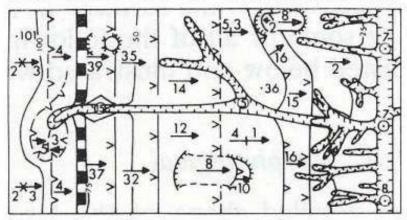


Fig. 2. Schematic diagrams illustrating the different principles of geomorphological map presentation for the same land surface. See paragraph 2.6.1, A-D.



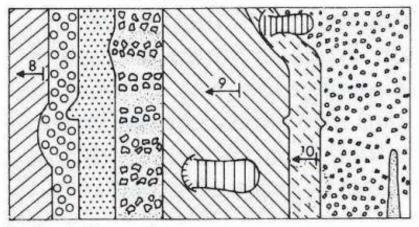
#### A. MORPHOLOGICAL /MORPHOMETRIC MAP



Morphological Mapping Symbols

- V V Convex break of slope
- A A Concave break of slope
- "V" Convex change of slope
- \_A\_A\_ Concave change of slope
- -10 ➤ Slope direction and angle
- Cliff > 45°
- Convex and concave breaks of slope in close association
- 4+2 Concave unit
- 2x3 ► Convex unit
- -73- Contours in metres
- -121 Spot height
- ⊙15 Depth of incision

#### B. MORPHOCHRONOLOGICAL MAP



Bedrock Succession

Planation Surface - Mid-Tertiary

OOOO Conglomerate

Sandstone

Mudstone (Highly weathered) Early-Middle Pleistocene

Retreating scarps pediment and

associated gravels

Siltstone - Late Pleistocene valley incision

Unconsolidated Sediments

River Terrace and infill – Devensian

River sand - recent

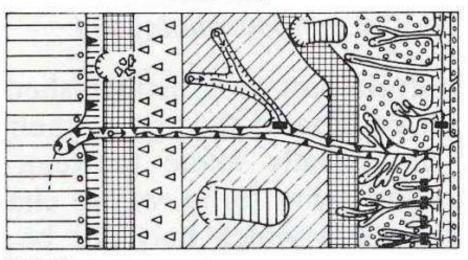
Angular boulders — intermixed — recent gravel & sand

Superficially Disturbed

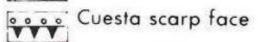
EIII Landslips — active

€10 Dip

#### C. MORPHOGRAPHIC MAP







TITITITI Rock wall

Scree – debris slope

/// Pediment

River terrace - and valley infill

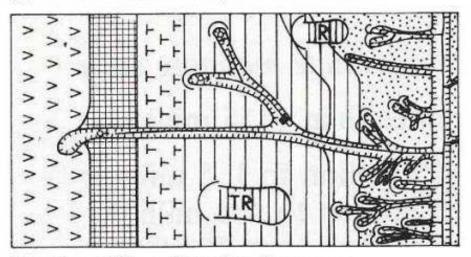
Incised valley side slope

(III) Landslides

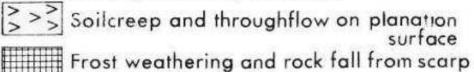
→ Spring → Waterfall — Permanent stream

Major gully IIIII Minor gully

#### D. MORPHOGENETIC/DYNAMIC MAP



#### **Dominant Slope Forming Processes**



Tosi wedinering did rock idii irom sedi

FF Talus creep on scree debris

Landslips on highly weathered mudstone

-active R=Rotational TR=Translational

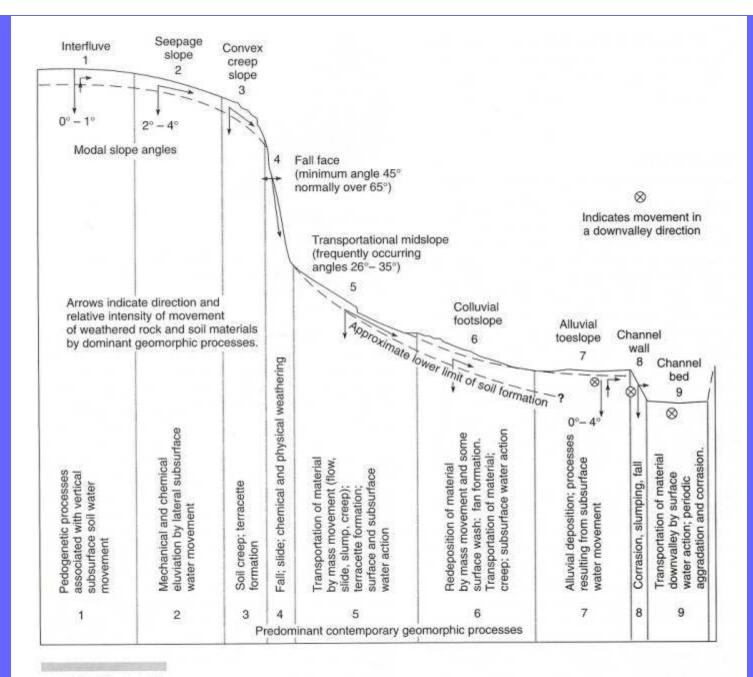
Potential instability on river terrace gravels

Wash on terrace

Gully erosion

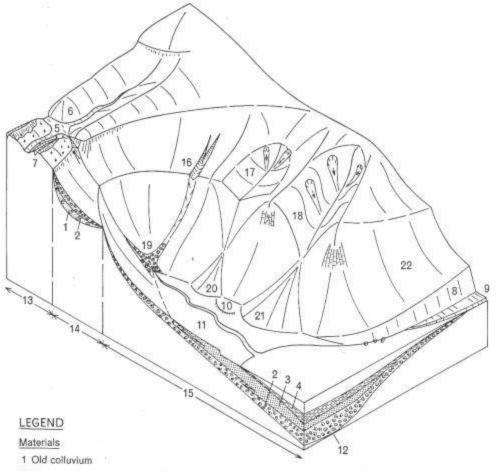
Actively eroding gully heads

→ Spring → Waterfall ← Permanent stream



#### FIGURE 4.45

Diagrammatic representation of the hypothetical nine-unit landsurface model. (Dalrymple 1968)



- 2 Young colluvium
- 3 Alluvium
- 4 Marine deposits

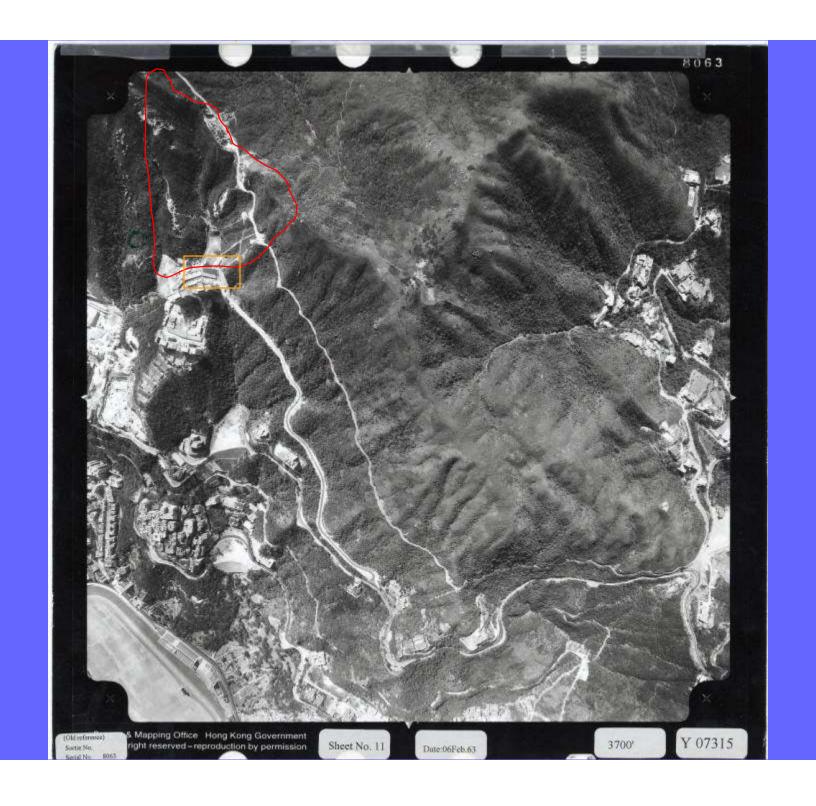
#### Landforms

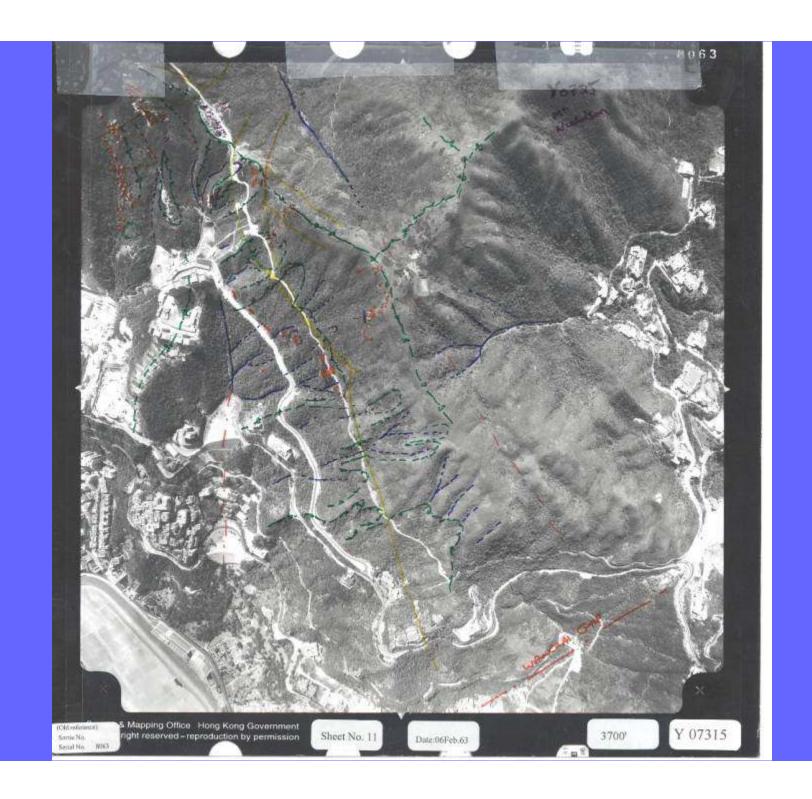
- 5 Upland valley
- 6 Deeply weathered hills
- 7 Ridge crest gully erosion
- 8 Coastal cliffs
- 9 Wave cut platform
- 10 Alluvial terrace
- 11 Floodplain
- 12 Submarine buried valley

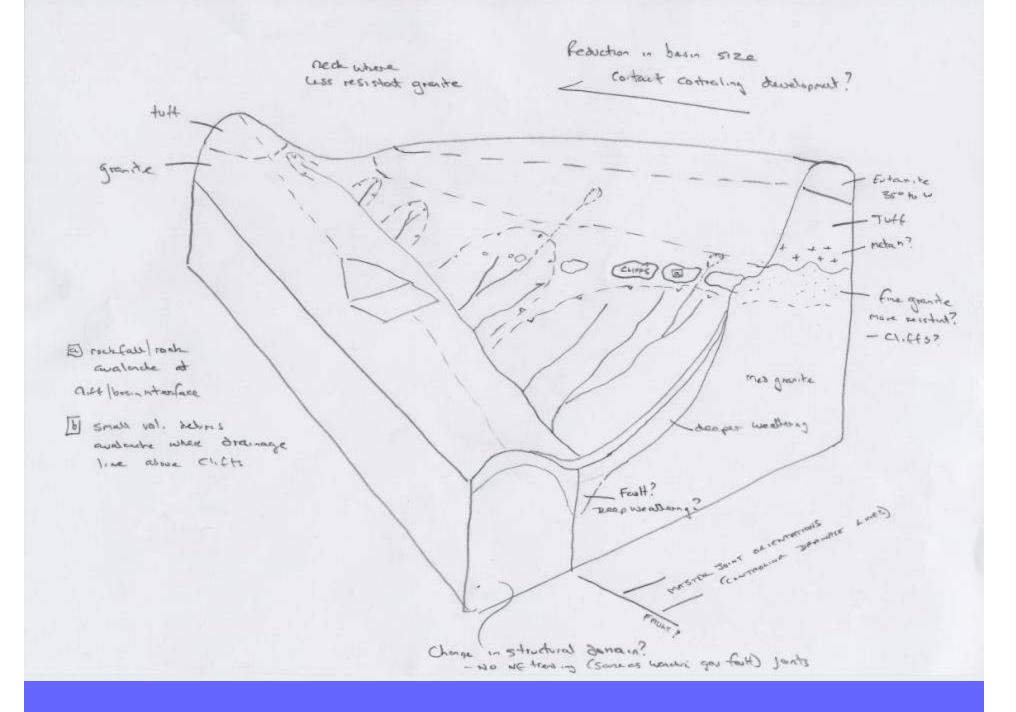
- Relict landforms on uplands
- 14 Older landforms
- 15 Younger landforms
- 16 Stream incising into superficial deposits
- 17 Initial incision has widened to a small valley, landslips active at margins
- 18 Deep, bowl-shaped valley between spurs, relict and recent instability on sides
- 19 Boulders in stream channel
- 20 Small colluvial fan
- 21 Large colluvial fan
- 22 Coastal slope (thin soils)

### Hansen, 1984

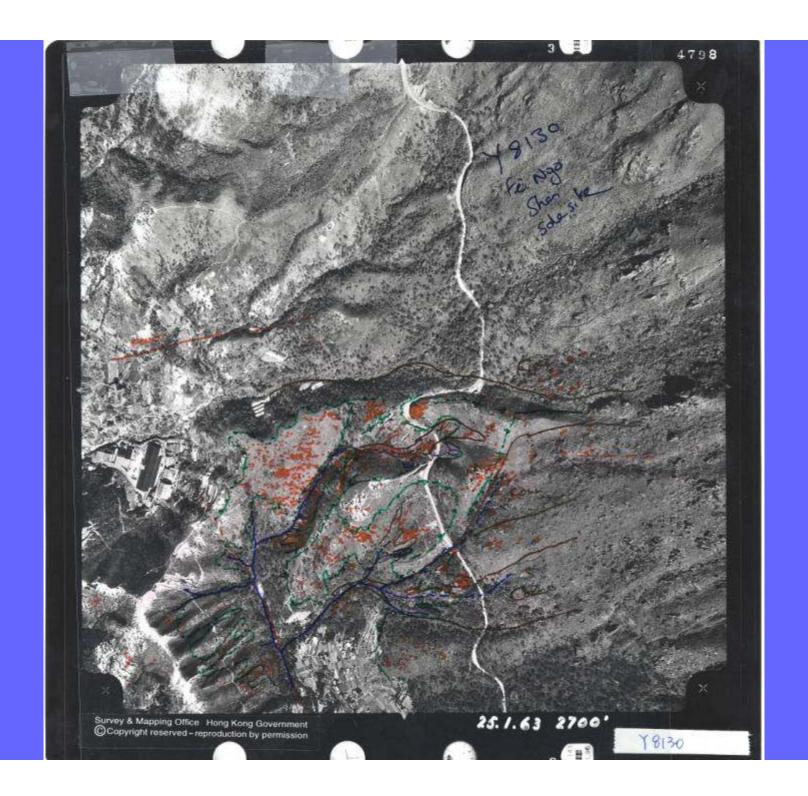
# Examples of developing geomorphological models

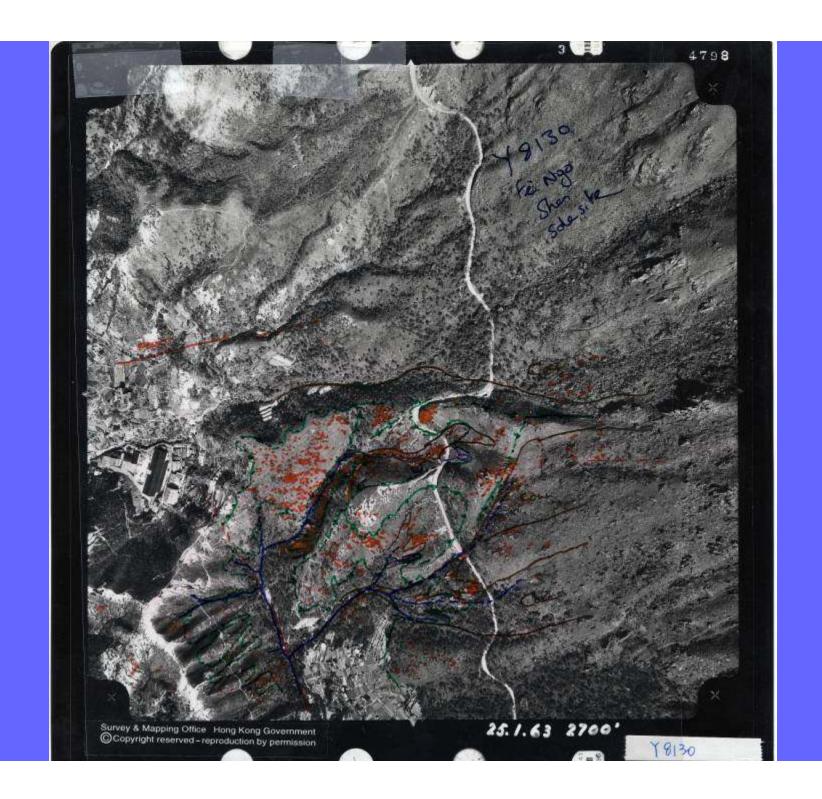


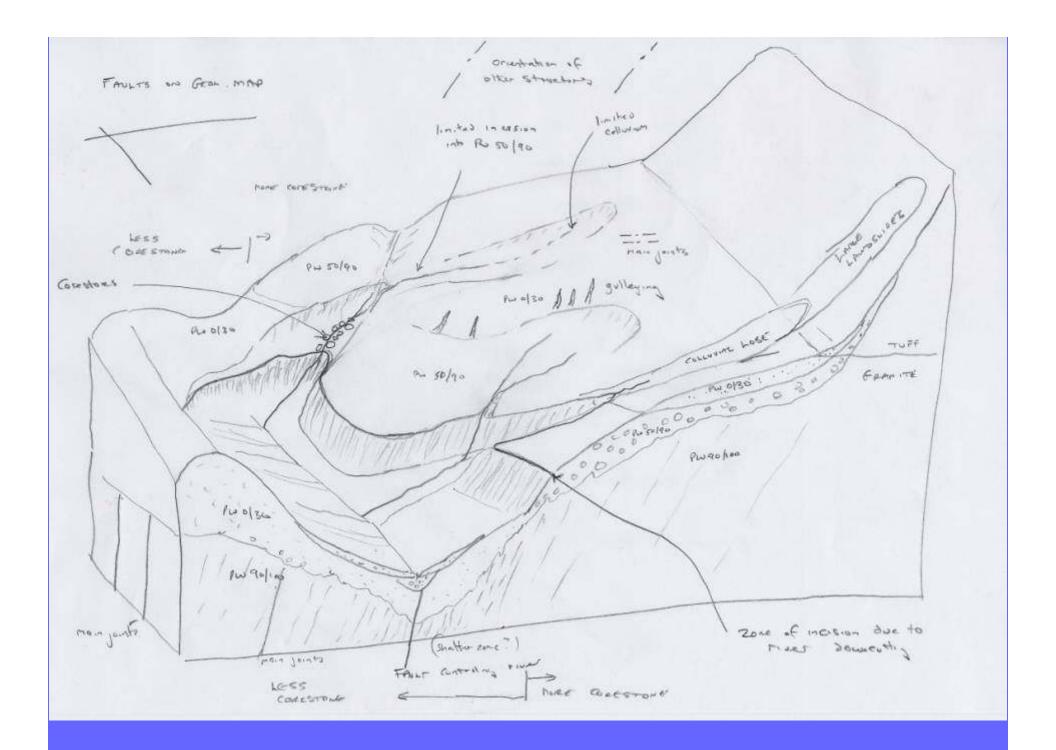


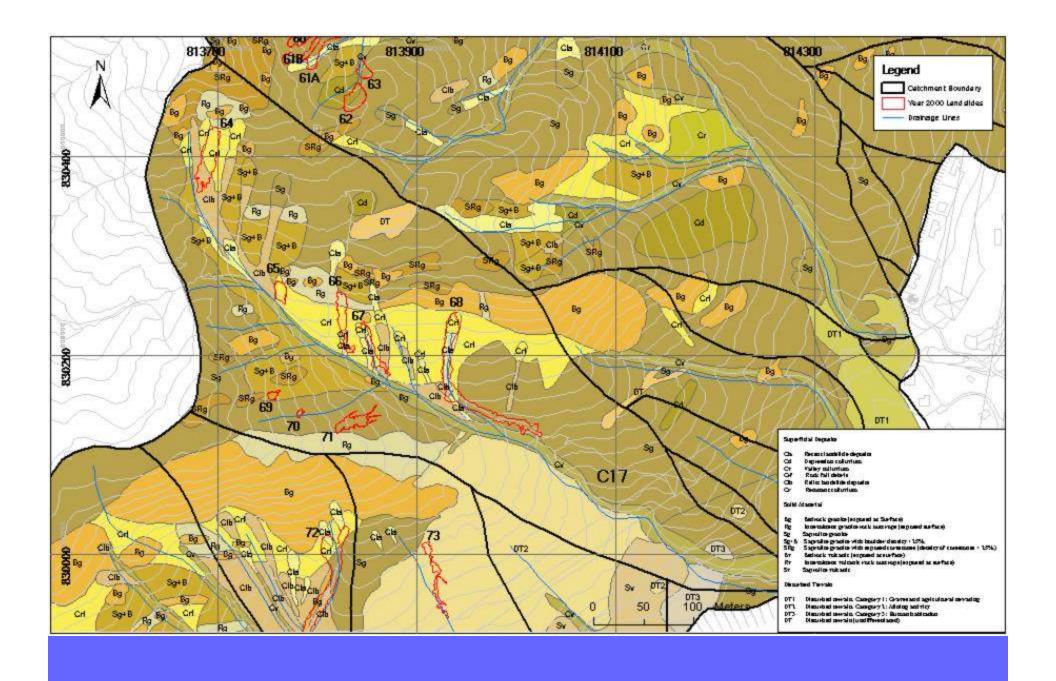


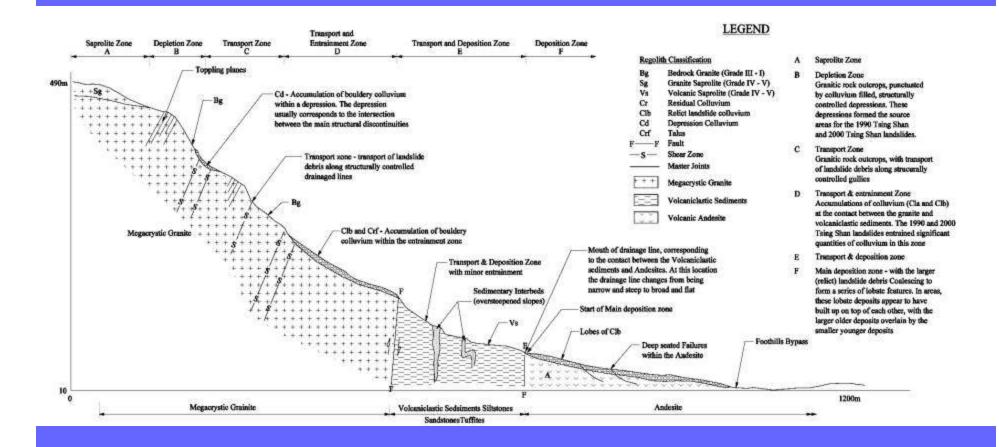












# Development of Hazard Models

Fletcher et al., 2002

#### Type I Landslide (Shallow)

Location: At the boundary between the upper and lower assemblage, generally at the heads of drainage lines.

Description: Shallow translational slide (~0.5 m to 2 m deep) typically along the old colluvium/saprolite interface of the upper landform.

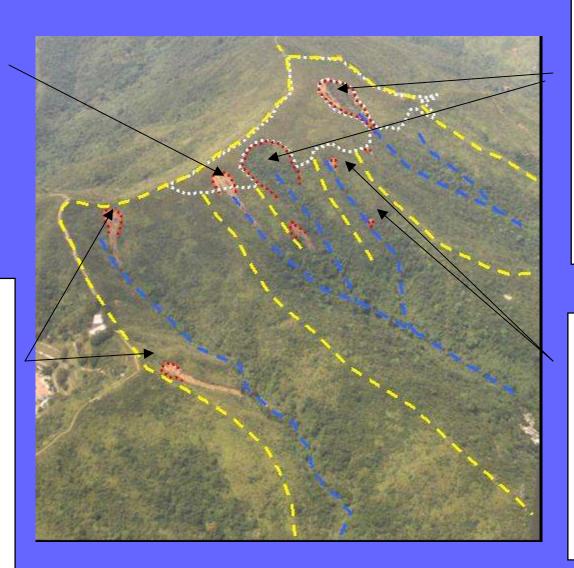
Typical Size: ~5 m to 25 m wide.

#### Type II Landslide

Location: At the head and flanks of incised valleys, generally located just below ridge lines within shallow hallows at the head of minor often poorly defined drainage lines. Typically located on the zero curvature line of a plan profile map.

Description: Shallow translational slide (~0.5 m to 2 m deep) generally along the colluvium/saprolite interface. Typically such landslides will occur episodically at the same location.

Typical Size: ~5 m to 15 m wide



# Type I Landslide (Large Multiple)

Location: At the boundary between the upper and lower landform assemblage, generally directly above the head of incising drainage lines.

Description: Generally large scale (~2 m to 10 m deep) within the deeply weathered saprolite forming the upper landform. Often retrogressive, resulting in multiple phases of landsliding.

Typical Size: ~20 m to 80 m wide.

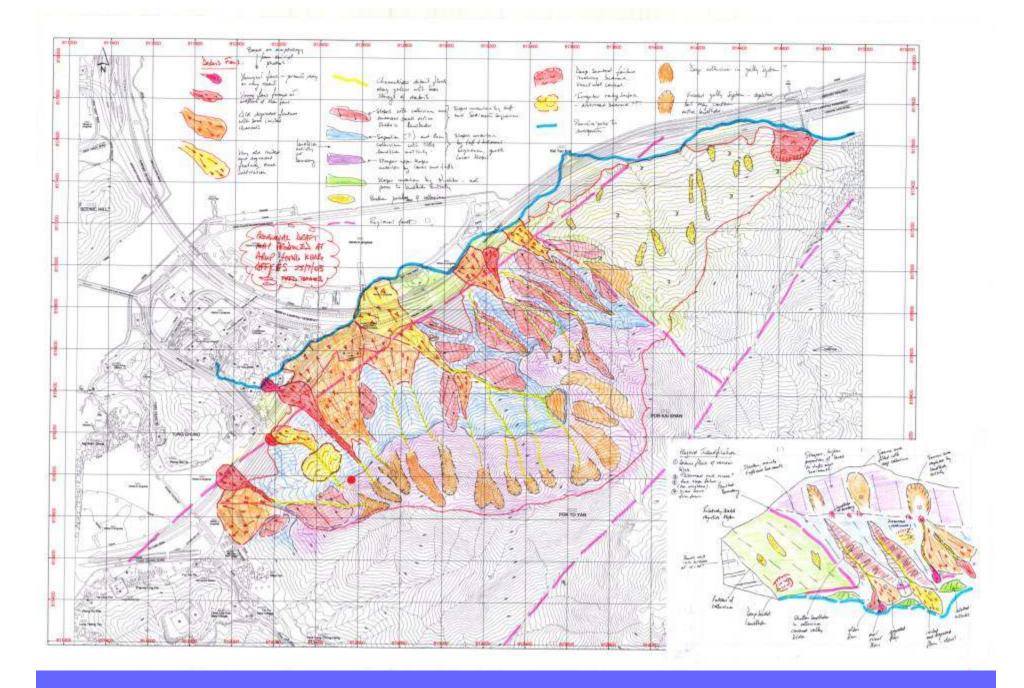
## Edge of Drainage Line Landslide

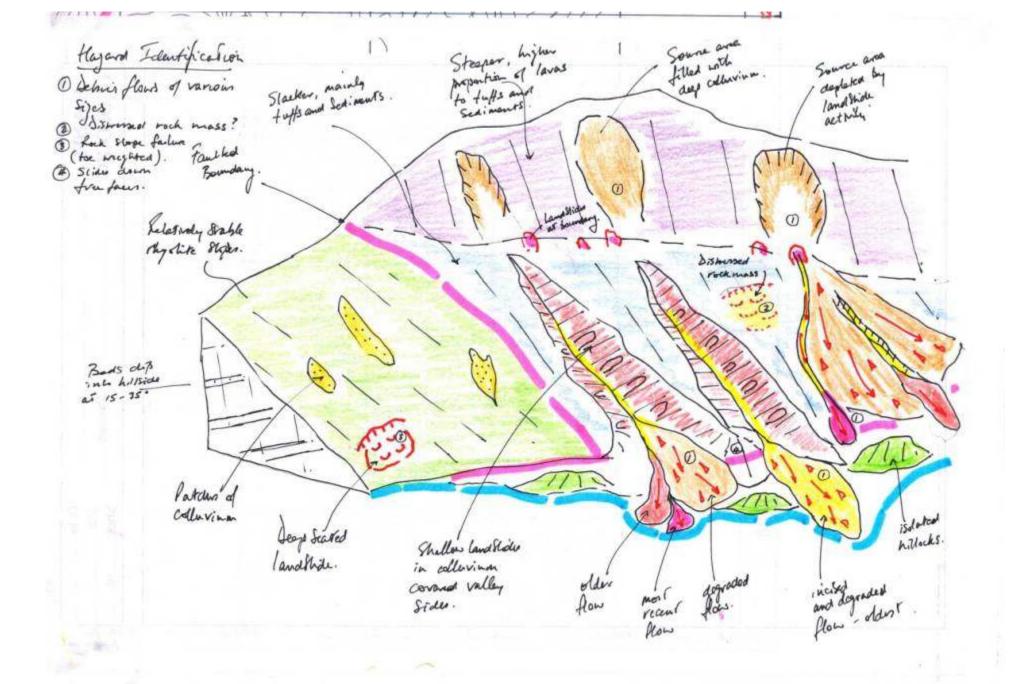
Location: Directly adjacent to eroding drainage lines, particularly at bends.

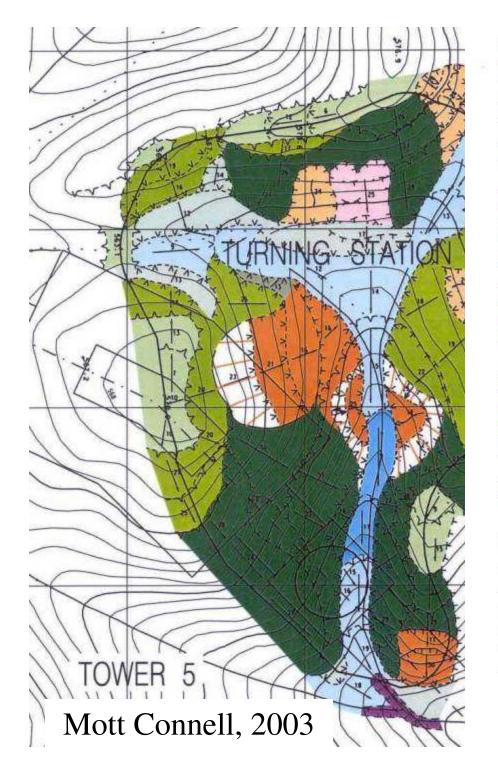
Description: Generally shallow rotational slumps (~0.5 m to 2 m deep) usually within colluvium deposits forming the flanks of the channel.

Typical Size: ~0.5 m to 5 m wide.

Halcrow, 2003







#### SLOPE FEATURES

RIDGES AND ISOLATED HILL TOPS, SUB-CONICAL HILLS WITH INTERCONNECTING RIDGES THAT ARE TYPICALLY NOT
GREATER THAN 15% WIDE. SLOPE ANGLES CAN BE WARLAGE RANGING UP TO 9" TO 15". COMSIDERED THAT BEDRICK IS
RELATIVELY CLOSE TO THE GROUND SUBFACE IS WITHIN 2% AND THAT SAPROLITIC SOILS DOMINATE THE MEAR SUBFACE.
SOME LOCAL SHEET EROSION IS IN EVIDENCE.

UPPER VALLEY SIDE SLOPES, TYPICALLY PRESENT DIRECTLY DOWNSLOPE OF THE RIDGES AND ISOLATED WILL TOPS.
SLOPES ARE USUALLY 16" TO 26". SUBFACE IS USUALLY LIGHTLY VEGETATED WITH GRASS. BOX.DERS TEND TO BE
ABSENT. A STIFF SAMOY SILT SOIL IS LOCALLY EXPOSED.

VALLEY SIDE SLOPES, TYPICALLLY PRESENT BETWEEN THE UPPER VALLEY SIDE SLOPES AND FOOTSLOPES OR

VALLEY SIDE SLIPES. TYPICALLY PRESENT BETWEEN THE UPPER VALLEY SIDE SLIPES AND FOOTSLOPES OR CHANNELS BELOW. SLOPES ARE USUALLY 28' TO 33'. ALTHOUGH LOCALITY AS SIGEP AS 35'. SOME BOULDERS UP TO THE ARE EVIDORY PARTICULARLY ON THE MORTHERN SIDE OF THE CATCHMENT. THAT ARE POORLY EMBEDGED. COMSIDERED THAT COLUMNAL OPPOSITS AND SAPROLITIC SOILS ARE RELATIVELY THIN ITS CYND. THESE SLOPES ARE CONSIDERED STABLE UNDER PRESENT CONDITIONS.

EQUISCOPES, DEVELOPED BELOW A CONCAVE CHANGE OF SLOPE AT THE BASE OF THE OTHER SLOPE UNITS. SLOPES RECORDED AS 12" TO 17".

#### CHANNEL FEATURES

CENTLE CHAMMEL BOTTOMS, COMPINED CHAMMELS THAT HAVE BASE WIDTHS UP TO TOM. SLOPE ANCLES WITHIN THE CHAMMELS ARE MODBALLY 9- TO 19". THERE IS LITTLE EVICENCE FOR EMPLAYERAL WAITER FLOW AND ACTIVE PROSIDE. COLLUVIAL ACCUMULATIONS WITHIN THE CENTRAL PART OF THE CATCHMENT ARE CONSIDERED TO BE UP TO 20" MICE.

CHANNELS, GENERALLY CONFINED CHANNELS THAT MAYE BASE WIDTHS TYPICALLY UP TO FOM. SLOPE ANGLES WITHIN THE CHANNELS ARE MORNALLY 12" TO 19". THERE IS SOME EVICENCE FOM EPHEMERAL MATER FLOW PARTICULARLY LOVER DOWN THE CATCHINGTURE WERE EPHEMERAL MATER SEPAMES WIGHT DOCUM.

STEEP CHARMELS, COMFINED CHARMELS THAT HAVE BASE WIDTHS TYPICALLY 15m. SLOPE ANGLES WITHIN THE CHARMELS ARE BORMALLY 319". THERE IS EXTENSIVE EVICANCE FOR EPHEMERIAL WATER FLOW AND ACTIVE ENDSIDM. THE CHARMEL SHOES OF UP 10 0.5m IS ALSO CYTOCHT.

#### INSTABILITY FEATURES

DESPATED DOTP TRANSLATIONAL LANDSLIDES, DEVELOPED DIRECTLY UPSLOPE OF THE CHANNELS AND CHINIE CHANNELS AND CHINIE CHANNELS OF STEEPER BACKSCAPPS TO 38" ABOVE SHALLOWER LANDSLIDE BODIES AT 16" TO 25". DEPTH OF FAILURE UP TO 24", WHERE THE TOE 15 ABSENT IT IS CONSIDERED THAT THIS MATERIAL HAS BEEN ENTRINED INTO THE CENTLE CHANNEL BOTTOMS COMMISTREAM. THERE IS NO EVIDENCE OF RECENT REACTIVATION.

DEGRADED SHALLOW TRANSLATIONAL LANDSLIDES, DISCRETE LANDSLIDE FEATURES THAT ARE TYPICALLY UP TO 30M-LONG AND 20M WIDE THAT WIGHT COMLESCE INTO LARGER AREAS. FAILURE DEPTH IS MOT CONSIDERED TO BE GREATER THAN THE BOOT OF THESE LANDIDES IS AT 207 - 32" AND APPEARS TO BE MARGINALLY STREEL EMDER PRESENT CONDITIONS. ASSOCIATED WITH PAST EROSION OF PREDOMINATELY SAPPELITIC SOILS.

HET IN SHEET PROSIDE AND SHALLOW TRANSLATIONAL LANGES TOING, AN AREA ZOW BY ZOW AFFECTED BY ACTIVE SHALLOW SOLL REMOVAL BY A COMBINATION OF PROCESSES. FAILURE DOPPH IS NOT CONSIDERED TO BE GREATER THAN THE SLOPE IN THIS AREA IS ZSY. ASSOCIATED WITH OMODING EROSION OF PREDOMINATELY SAPROLITIC SOILS SUCH THAT REPORCE CREOSURES AND PRESENT.

#### MORPHOLOGY

A A CONVEX BREAKS OF SLOPE

V V CONCAVE BREAKS OF SLOPE

A CONVEX CHANGES OF SLOPE

\_V\_ \_V. CONCAVE CHANGES OF SLOPE

ALLILL BREAKS OF SLOPE

THE CHANCES OF SLOPE

SLOPE DIRECTION WITH AMOLE IN DECREES

UNDALATING SLOPE DIRECTION WITH ANGLE IN DEGREES

#### TOWER 5 CATCHMENT GEOMORPHOLOGICAL MODEL

SUB-ROUNDED HILL TOP ISOLATED HILL TOP AREA OF SHALLOW EROSION DRY CHANNEL & TRANSLATIONAL SLIDING SADDLE UPPER CHANNEL ELONGATE RIDGE & SLOPES GRASSED OLD TERRACETTES DECOMPOSED TUFF BEDROCK EXPOSED IN ERODING AREA RELATIVELY WIDELY SPACED MASTER JOINT OLD DEGRADED SETS IN TUFF BELOW TRANSLATIONAL LANDSLIDE ISOLATED HILLS LIMITED RESIDUAL SOIL DEVELOPMENT ON DEGRADED OLD SHALLOW RIDGE TOPS TRANSLATIONAL LANDSLIDE ORIENTATION OF RIDGE LINES PARALLEL TO MAJOR SUB-VERTICAL WEATHERED MANTLE WITH SILTY SOILS & JOINT SETS CORESTONES JOINT SETS RELATIVELY WIDELY SPACED BELOW POTENTIALLY UNSTABLE RIDGE LINES BOULDERS ON GROUND SURFACE TYPICAL TUFF BEDROCK BENCH DEVELOPED WHERE WEATHERING MORE INTENSE LOCAL CORESTONES DEVELOPED IN OLD FAILURE AT CHANNEL SLIGHTLY & CONFLUENCE MODERATELY WEATHERED TUFF PARTIALLY EMBEDDED BOULDERS **APPROXIMATE** JOINT SETS MORE TOWER FOOTPRINT CLOSELY SPACED SUB ROUNDED BOULDERS BENEATH CHANNEL & LOCAL BEDROCK EXPOSED IN CHANNEL BED SMALL TREES & SHRUBS IN LOWER CHANNEL LOCAL EVIDENCE OF EROSION CHANNEL WITH INCISION

FIGURE 2

# Advantages of a geomorphological approach

# Assists in determining:

- the type of natural terrain hazards present
- the location of potential source areas
- the estimation of source volumes for hazard types
- the likely frequency of the hazard types
- the presence of entrainable material

# Physical Investigations of Natural Terrain



#### Natural Terrain Hazard Study for Tsing Shan Foothill Area Landslide Field Proforma

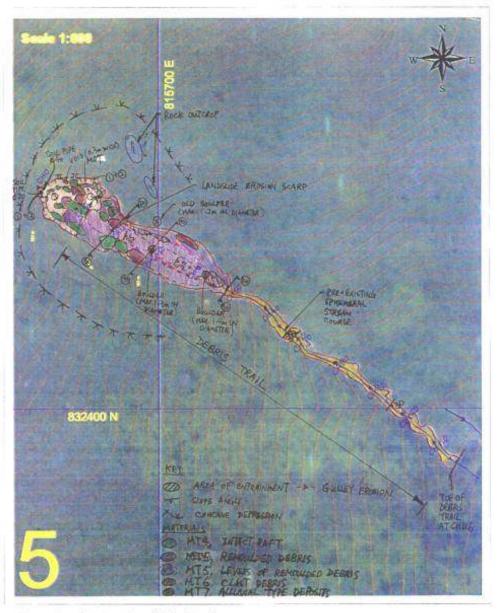


Figure 44 Sketch plan of debris trail

(Note: based on orthorectified photograph with material type average thicknesses and maximum clast diameter for each deposition area)

#### Natural Terrain Hazard Study for Tsing Shan Foothill Area Landslide Field Proforma

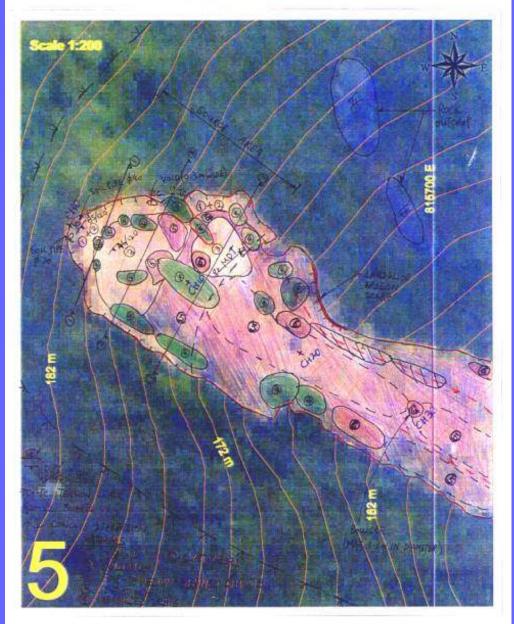


Figure 2 Sketch plan of landslide source

(Note: based on orthorectified photograph with average thicknesses of depletion on overlay).





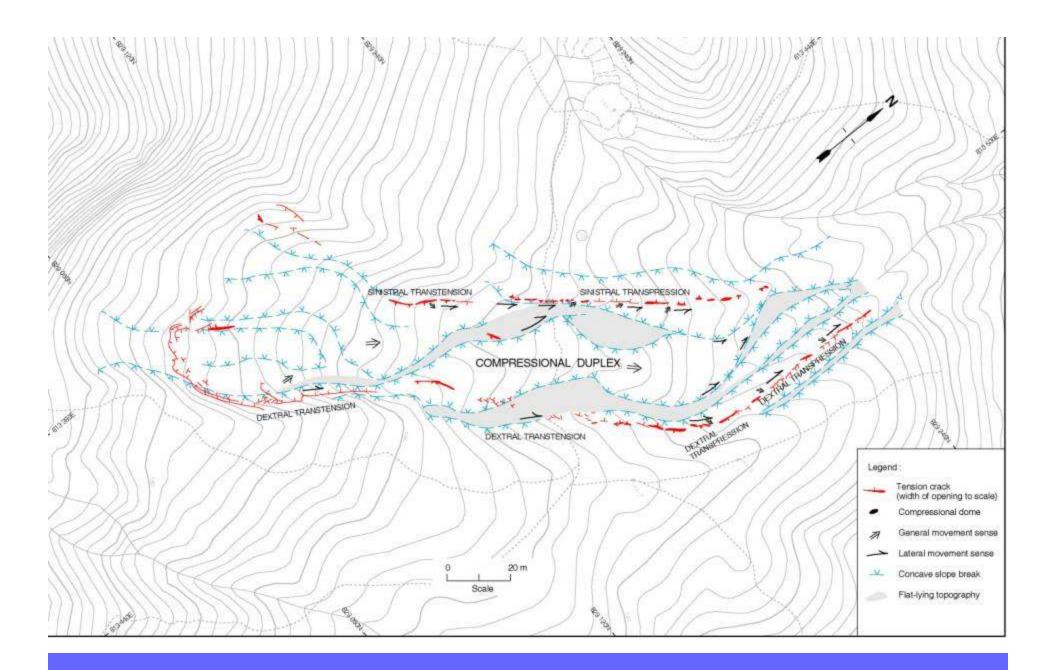












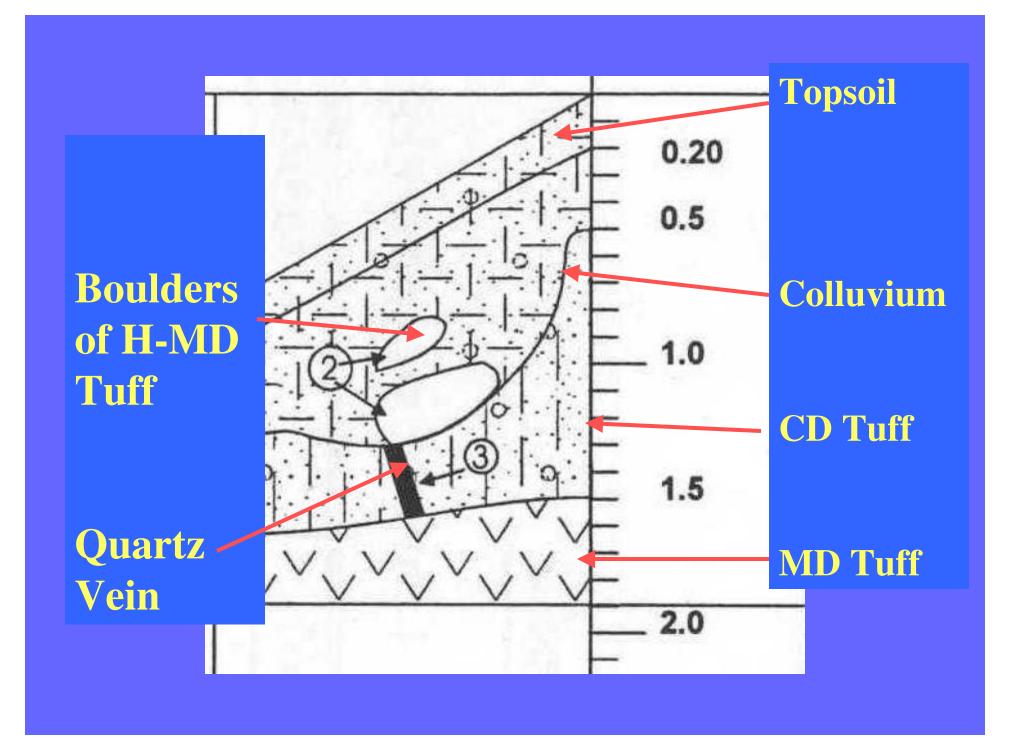
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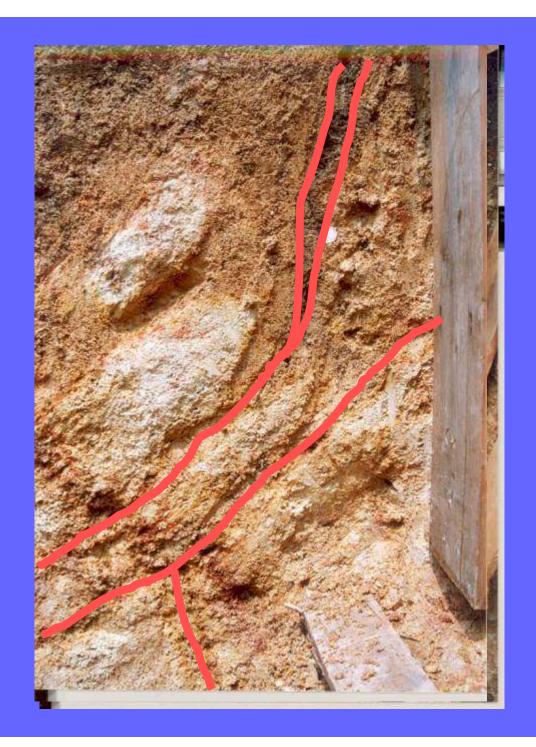
## GEOTECHNICS & CONCRETE ENGG. (H.K.) LTD. GROUND INVESTIGATION DEPARTMENT

HOLE NO.

BHC2

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FLUSHING MEDIUM Air-Foam								ORIENTATION Vertical					GROUND LEVEL 134.29 mPG		
Drilling Progress	Casing size	Water level (m) & Time	Total core Recovery %	Solid core Recovery %	R.Q.D.	Fracture		Samples	Reduced	Depth (m)	Legend	Grade		Description	
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Colluvium clayey silt

**CD Tuff** 

**Tension Crack** 

Colluvium

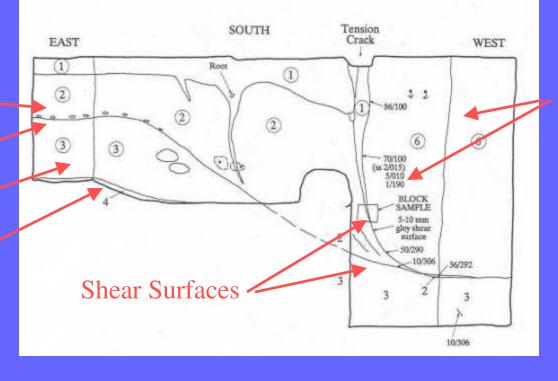
clayey silt with some gravel cobbles & boulders

Old Colluvium clayey silt

Stone line

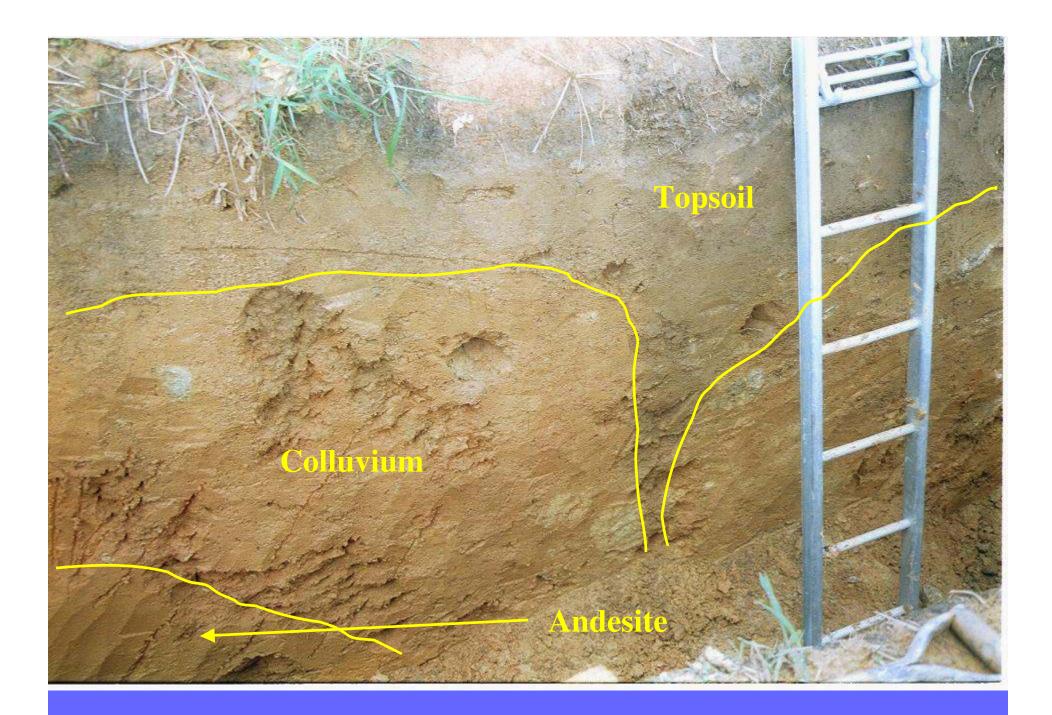
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## Landslide Debris

clayey sand with much gravel cobbles & boulders



## Thank You